

碳酸盐岩微孔隙的分类、成因及对岩石物理性质的影响

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摘要:碳酸盐岩微孔隙(简称微孔)是孔径小于50 μm 的微米-纳米级孔隙,目前有多种单因素和综合分类方案,以Kaczmarek的综合划分方案应用较为广泛,该方案将泥晶形态和物性相结合。微孔的表征方法主要有流体侵入法、图像分析法和数学统计法,并通过定性-定量组合实现对微孔的综合表征。微孔主要发育于泥晶-粉晶之间,原生沉积是微孔发育的基础,成岩作用是微孔形成的关键,重要的成岩作用包括矿物的转化作用、白云石化作用以及重结晶作用等。而泥晶体形态和排列方式决定微孔的空间几何结构,从而影响岩石的物理性质。随着对碳酸盐岩微孔认识的不断加深,需要进一步结合中国实例进行理论完善(包括微孔的分类、成因等),同时应用更先进的高精度微区测年、激光扫描共聚焦显微镜和数字岩心模型等表征技术,建立和完善微孔碳酸盐岩储层岩评价标准,以对中国大量发育的深层碳酸盐岩储层,尤其是其微观特征的研究提供理论和技术支撑。

关键词:表征方法;成因机理;物性特征;成岩作用;沉积相;微孔;碳酸盐岩储层

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Classification and origin of micropores in carbonates and their effects on physical properties of rocks

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Abstract: Micropores in carbonate rocks are micron- to nano-sized pores with pore sizes less than 50 μm . Various single-factor and comprehensive classification methods for these micropores are available, with Kaczmarek's comprehensive scheme being widely applied due to its integration of aphanocrystalline morphologies and physical properties. While the fluid intrusion method, image analysis, and mathematical statistics serve as primary characterization methods for micropores, a combination of qualitative and quantitative methods can achieve their comprehensive characterization. Micropores are primarily found between aphanocrystallines or very fine crystallines, primarily growing on primary minerals. Diagenetic processes, especially mineral transformations, dolomitization, and recrystallization, significantly influence micropore formation. Additionally, the morphologies and configurations of aphanocrystallines govern the spatial geometry of micropores, further influencing the physical properties of rocks. As the understanding of micropores in carbonates deepens, it is necessary to refine existing theories about micropore classification and origin by combining

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the context of China. Furthermore, there is a need to establish and refine the evaluation criteria for rocks of microporous carbonate reservoirs using more advanced characterization techniques such as in-situ microanalytical methods, confocal laser scanning microscopy (CLSM), and digital core modeling. This will be of theoretical and technical support to research on the micro characteristics of deep carbonate reservoirs prevalent in China.

Key words: characterization method, genetic mechanism, physical property, diagenesis, sedimentary facies, micropore, carbonate reservoir

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碳酸盐岩微孔隙(简称微孔)是孔径范围属于微米–纳米级的孔隙,既可以发育在石灰岩的基质或胶结物之间、颗粒内^[1–4],也常见于微生物灰岩/白云岩之中^[5–7]。微孔可以从孔径大小及岩石物理特征两个方面界定。就孔径大小而言,不同学者对微孔最大孔径值有多种界定,范围在2 nm至100 μm 不等(图1);也有以孔喉半径来划分,小的在0.3 μm ^[8],大的可达50 μm ^[9]。不同研究目的和技术方法造成了微孔孔径界限的差异。例如,在光学显微镜下,Archie认为微孔是不可见孔隙,直径小于10.0 μm ^[10];基于多种类型微孔的平均孔径,Choquette和Pray认为微孔直径小于62.5 μm ^[11];在透射电镜和扫描电镜下,Pittman认为至少在一个维度上,微孔直径小于1.0 μm ^[12]。从岩石物理尤其是油气采收率的角度,Fullmer等划分微孔与宏观孔隙的界限是10.0 μm ^[13]。随着环境扫描电镜、纳

米CT扫描仪等高分辨率设备的应用,纳米级孔隙也被纳入到微孔的研究范围中^[7, 14–15]。

常见碳酸盐岩微孔孔径上限为10 μm ^[16–22],而不同的微孔类型会导致孔径上限不同^[23],在研究中也常见几十微米的微孔。本文考虑微孔上限的普适性并参考Arve Lønøy的碳酸盐岩孔径分类方案^[23],将微孔的上限定为50 μm 。

在国外多个盆地不同埋深和层系的微孔碳酸盐岩(microporous carbonates)储集层中均有工业油气的发现^[7, 17, 19–20, 24–29]。与公认的礁滩型储层、岩溶储层以及白云岩储层不同,微孔碳酸盐岩储层的微孔率(微孔孔隙度/总孔隙度)大于90 %^[21],属于非常规储集岩。微孔不仅发育在海相泥晶灰岩/泥晶白云岩和海相颗粒灰岩/颗粒白云岩中,也发育在湖相微孔碳酸盐岩中。中国微孔碳酸盐岩长期未得到重视,近年来在柴达木

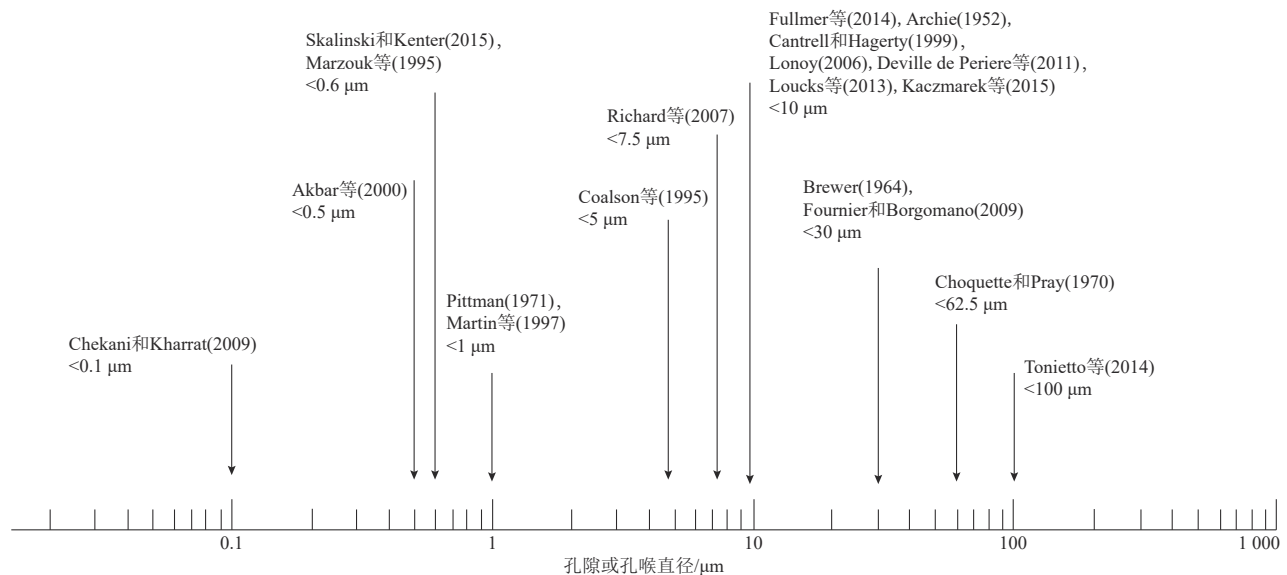


图1 不同学者对微孔直径的定义(据参考文献[1]修改)

Fig. 1 Delineation of micropore diameters provided by different researchers (modified after reference [1])

(横轴是以微米表示孔隙或孔喉直径的对数刻度;箭头表示不同作者使用的孔径大小上限。)

表1 中国碳酸盐岩微孔研究现状^[21, 31-38]Table 1 Current status of China's research on micropores in carbonate rocks^[21, 31-38]

地区层位	沉积岩相	孔径范围	成因
塔里木盆地顺南地区奥陶系一间房组	生物黏结灰岩、藻屑灰岩	10 ~ 20 μm 到亚微米尺度	有机质演化与重结晶作用共同控制、热液流体改造、早期烃类充注
四川盆地川东地区二叠系茅口组	缓坡型碳酸盐岩台地环境, 泥灰岩	1 ~ 1 000 nm	海泡石发生脱水和硅质的析出, 晶体体积变小; 矿物颗粒内部溶蚀孔, 主要为纳米级孔隙
塔里木盆地柯坪露头区震旦系奇格布拉克组	微生物白云岩	<10 μm	微生物热解生酸溶蚀作用
塔里木顺南地区中-下奥陶系	藻灰岩、藻云岩、非生物灰岩、白云岩	<10 μm	准同生期矿物转化、Ostwald熟化、热液溶蚀
羌塘盆地二叠系龙格组	微晶灰岩及生物屑灰岩、白云岩	半径<20 μm	白云石化作用、大气淡水溶蚀作用、压溶作用
四川盆地川西北地区二叠系栖霞组	泥晶颗粒云岩	7.3 μm	
柴达木盆地英西地区古近系干柴沟组(湖相)	泥晶云岩	一般<1 μm	白云石化作用, 有机酸溶蚀; 大气淡水选择性溶蚀
粤中地区三水盆地古近系布心组(湖相)	泥晶灰岩内部, 部分生物灰岩	孔隙平均半径1.799 μm	基质孔和少量溶蚀作用
鄂尔多斯盆地大牛地中奥陶统马家沟组	泥晶-粉晶白云岩	粉晶白云岩中的晶间孔20 ~ 30 μm ; 而微晶白云岩中的晶间孔一般<10 μm	白云石化作用

盆地古近系^[30]、塔里木盆地震旦系和奥陶系^[31-32]、四川盆地二叠系^[33]等均有微孔的报告(表1), 最近在四川盆地茅一段的微孔泥灰岩(微孔孔径仅为4 ~ 1 400 nm)中取得了天然气勘探的突破^[33], 展现了非常规(微孔)碳酸盐岩储层的巨大潜力。中国西部碳酸盐岩油气储集层普遍具有孔隙度低、埋深大的特点, 系统开展碳酸盐岩微孔的表征、成因和成储潜力等研究, 不仅可以深化碳酸盐岩微孔形成机理的地质认识, 还可为中国西部超深层碳酸盐岩储层研究提供新的思路, 具有重要的理论和应用价值。

1 微孔分类及表征方法

1.1 单一因素分类

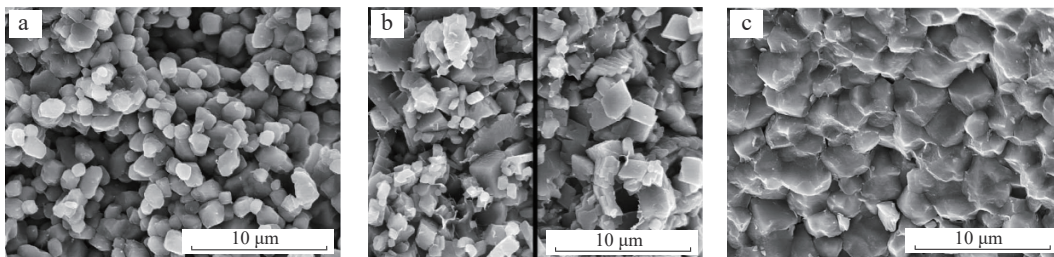
依据微孔与组分的关系, 如基质内、颗粒内、纤维状胶结物之间、等粒状胶结物之间, 可分为基质孔、粒内孔和晶间孔等^[1, 39-40]。而依据孔径大小可分为微米级和纳米级孔隙^[7, 41], 甚至细分为极细孔(0.01 ~ 2.00 μm)、细孔(2.00 ~ 4.00 μm)、中孔(4.00 ~ 6.00 μm)、粗孔(6.00 ~ 8.00 μm)和极粗孔(8.00 ~ 10.00 μm)。也可以根据孔隙的成因, 将微孔分为晶间孔^[4, 16, 19, 21]、晶间溶蚀孔、生物体腔孔和有机孔(有机质热演化形成的沥青收缩孔)等^[42-43]。

1.2 综合分类

定量的综合分类方案主要依据的是泥晶矿物晶体结构特征及实验室内获得的岩石物理参数值, 以Kaczmarek的划分方案应用较广泛^[19]。Kaczmarek对全球显生宙的灰岩储层进行了研究, 包括对矿物学、地球化学及岩石物性实验数据等方面的分析, 将微孔划分为I型、II型和III型。I型发育粒状半自形晶体结构(granular-subhedral), II型发育粒状自形(granular-euhedral)、疏松簇状(clustered-loose)和紧密簇状(clustered-fused), III型发育部分接合状(fitted-partial)和紧密接合状(fitted-fused)晶体结构。3种微孔的各项岩石物性参数依次降低: 孔隙度分别为>20%, 10% ~ 20%和<10%, 渗透率分别为 $(1.0 \sim 20.0) \times 10^{-3}$, $(0.1 \sim 1.0) \times 10^{-3}$ 和 $<0.1 \times 10^{-3} \mu\text{m}^2$, 平均喉道半径分别为0.70, 0.20和0.06 μm (图2)。中国也有类似的分类, 卫端(2020)在对塔里木盆地顺南地区鹰山组和一间房组的微孔灰岩进行研究时^[44], 在岩性和泥晶结构特征的基础上, 结合岩石核磁共振 T_2 谱分布(核磁共振弛豫成像)和岩石物理性质对微孔进行分类。

1.3 表征方法

碳酸盐岩微孔的表征, 对微孔碳酸盐岩的分类和流体渗流机理等具有指导意义。但针对碳酸盐岩微孔

图2 微孔碳酸盐岩储层3种微孔类型的晶体形态^[19]Fig. 2 Crystalline morphologies of three major types of micropores in microporous carbonate reservoirs^[19]

a. I型,粒状半自形晶体结构;b. II型,粒状自形晶体结构和簇状晶体结构;c. III型,接合状晶体结构

的表征方法有限,相对于宏孔研究较薄弱。电子计算机断层扫描(CT)、核磁共振(NMR)、激光扫描共聚焦显微镜(CLSM)、场发射扫描电子显微镜(FESEM)、数字图像分析(DIA)、三维纳米孔建模、孔隙尺度成像和建模的深度学习等仪器和技术的应用,为微孔的研究提供了方法支撑^[1, 41, 45-53]。部分页岩和致密砂岩的微孔表征技术方法可以用在碳酸盐岩微孔的表征上^[21]。碳酸盐岩微孔表征的技术方法根据原理分为3类:基于流体侵入法(高压压汞法 HPMT、气体吸附法、混相驱替法^[54]等)及核磁共振法等分析孔隙及孔道特征的定量表征;基于各类图像分析法(光学显微镜、扫描电镜等)、CT扫描法等对孔隙形态进行描述的定性表征;以分形维数为代表的数学统计。

目前,微孔的表征更趋向于定性和定量的综合表征,即用多个方法进行相互验证,从多个方面对微孔进行表征^[21, 55-57]。例如在塔里木盆地顺南地区中-下奥陶统深层微孔碳酸盐岩的研究中,用扫描电镜法、核磁共振法、高压压汞法以及分形维数统计法全方面对微孔进行综合表征^[21]。同时,针对微孔的表征方法在非均质碳酸盐岩的孔隙表征中具有局限性,因此需要复合型孔隙的表征方法,实现对宏观孔和微观孔进行分区-分层表征^[57]。

而在实际生产应用中,更多关注的是微孔的物性特征。传统的研究以实测为主,以柴达木盆地英西地区湖相微孔白云岩的研究为例,采用包括常规压汞及扫描电镜确定孔喉分布范围,核磁共振离心联测可动孔喉下限及物性下限,用压汞曲线法和分布函数法进行验证,综合确定微孔白云岩储层的物性特征^[56]。同时,还可以利用两种表征方法的互补性对微孔物性特征进行研究,例如通过核磁共振法和高压压汞法曲线相减的方法,低成本和高效地得到微孔中连通孔隙和不连通孔隙的分布特征^[34, 58]。在计算机技术的支持下,以孔隙连通性为核心,提出三维成像、数字岩心,连通域形态学分析和连通域数字模型提取等研究技术和

方法^[59-60]。这类研究方法能够更好地反映微米-纳米级孔隙空间的本质特点,兼顾微米-纳米级孔隙的动态和静态特性,但这些技术受限于其代表性、分辨率以及成本。

2 微孔成因与典型发育模式

微孔广泛分布在碳酸盐岩颗粒内部、基质以及胶结物中,赋存于泥晶-粉晶之间^[1, 4, 16, 19, 21, 61](图3),因此前人在研究碳酸盐岩微孔成因时更关注原生矿物学特征的变化、沉积相的影响以及成岩作用的改造。

2.1 原生沉积物与沉积条件

古老碳酸盐岩中几乎所有的微孔都发育在泥晶低镁方解石和泥晶白云石中,与现代碳酸盐沉积物相比,在矿物成分和结构等方面均有所差异^[1, 4, 18, 62-66]。因此,学者们大量讨论了原生碳酸盐沉积物的结构和矿物成分对微孔发育的控制作用^[2, 16, 20, 67]。除此之外,沉积条件对微孔的形成也有一定的影响。

2.1.1 原生矿物成分决定微孔形成的潜力

在文石海和方解石海的不同时期,沉积物矿物成分差异影响了微孔的发育,古老微孔碳酸盐岩的原生矿物成分是低镁方解石还是文石/高镁方解石,学界尚无统一论。本文认为,不论是低镁方解石还是高镁方解石/文石,都可以形成微孔,但高镁方解石和文石更有利于微孔的产生,而矿物成分差异影响微孔发育的潜能和成岩过程。

微孔灰岩多发育于白垩纪方解石海时期以及新生代方解石海向文石海过渡时期^[28],特别是中东地区白垩系发育大量微孔灰岩,这表明低镁方解石有利于形成微孔。通过对其机理展开研究,有学者认为文石和高镁方解石在向低镁方解石转化的过程中,体积增加,孔隙度减少,原生结构遭到破坏,从而不利于微孔的保

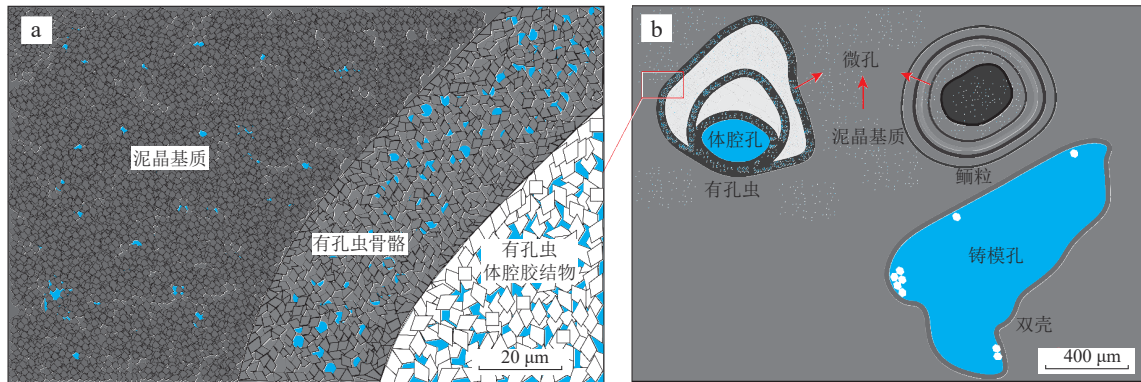


图3 不同尺度下各组分中的微孔

Fig. 3 Micropores in mineral components on different scales

a. 微米尺度的微孔; b. 毫米尺度的微孔(“蓝雾”)

存^[22, 68], 而低镁方解石会在 Ostwald 熟化 (Ostwald ripening) 影响下发生重结晶作用, 使得低镁方解石泥晶排列更加均一, 有利于晶间微孔的形成^[18, 69]。

然而在文石海时期, 同样能够发育微孔碳酸盐岩 (石炭系、新近系)^[1, 12, 65, 70-71]。通过实验证明, 文石向方解石的转化对孔隙度的影响很小^[72], 更多人认为从亚稳的文石和高镁方解石转化为低镁方解石这个过程^[2, 16, 20, 73-75], 对于微孔的形成来说是一个建设性的过程 (孔隙不是在这个过程中产生的, 而是从原始沉积物中继承的)。事实上, 无论是在基质还是颗粒中, 高镁方解石和文石都表现得更容易发育微孔, 特别是在高镁方解石的颗粒中更有利于微孔的形成^[16]。例如, 在德克萨斯州 Pawnee 油田中, 文石质颗粒 (双壳类) 被溶解形成铸模微孔, 而 HMC 颗粒 (包括有孔虫、红藻以及其他类似颗粒) 形成了晶间微孔^[20]。而且高镁方解石向低镁方解石转化为孔隙流体提供大量 Mg^{2+} , 当 Mg^{2+} 过剩时将导致白云石化, 有利于形成白云石晶间微孔。反应机制表明, 当文石转化为白云石时体积缩小 6.15%, 而方解石转化为白云石时体积缩小 14.81%^[76]。另外值得一提的是, 原生矿物成分控制成核位点的数量, 对泥晶尺寸起到控制作用^[64, 72, 77-78], 从而影响微孔的形态。高镁方解石相较于文石, 具有更多成核位点, 形成的低镁方解石晶体较细, 更有利于微孔的形成^[64, 72, 78]。

2.1.2 开放的原生微观结构利于微孔形成

当颗粒内部的微观结构为泥晶-粉晶结构或是有利于液体通过的开放结构时, 均有利于微孔的形成。在全球显生宙微孔灰岩统计中, 发现似球粒、鲕粒、珊瑚、有孔虫和藻类等颗粒更容易发育微孔^[19]。Walter 和 Morse 通过实验证明, 在某些化学条件下, 颗粒显微

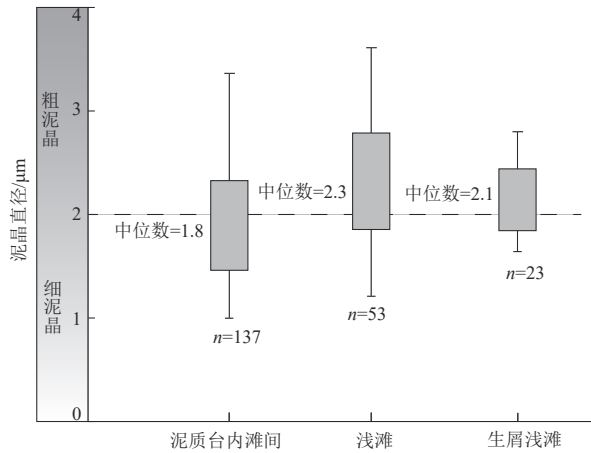
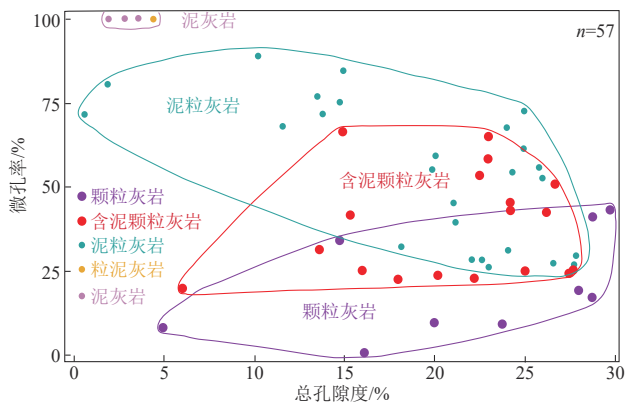
结构可以超越矿物溶解的热力学约束^[79-80]。在实验室条件下, 在低镁方解石欠饱和和溶液中, 具有泥晶-粉晶结构的文石颗粒溶解速度更快^[80-81]。同样, 开放的骨骼结构有利于液体通过, 也更容易形成微孔^[2]。

2.1.3 沉积环境影响泥晶发育及微孔丰度

沉积条件的差异影响泥晶的形态, 不同沉积环境中碳酸盐岩的微孔率也有所不同。①沉积条件控制泥晶发育程度, 进而影响微孔的形态。沉积环境水体能量越高, 结晶程度越高。Matthieu 对中东地区 3 个白垩纪微孔碳酸盐岩的沉积环境进行统计, 发现沉积水动力条件与微孔分布有一定相关关系, 高能浅滩相更多发育的是粗泥晶 (泥晶直径 $>2\ \mu\text{m}$) 晶间孔, 而低能泥质内台地相中主要发育细泥晶晶间孔^[17] (图 4)。现代碳酸盐岩沉积物同样表明了泥晶结晶程度与水动力强度具有函数关系^[82-83], 但是在古老的碳酸盐岩中, 由于受到成岩作用的改造, 而很难精确描述这种关系^[17, 84]。②沉积条件与微孔丰度具有相关性, 伴随水动力减弱, 灰泥含量增加, 总孔隙度减少, 微孔的总比例上升 (图 5)。有学者对微孔碳酸盐岩储层提出严格定义, 即微孔占总孔隙 90%、甚至 95% 的碳酸盐岩储层为微孔碳酸盐岩储层^[21, 85]。因此认为微孔碳酸盐岩储层主要分布在大孔隙发育少、灰泥含量较高的低能沉积相中^[86-87]。

2.2 成岩作用

在微孔形成的过程中, 成岩作用是必要的改造条件, 包括了早期成岩过程的改造和埋藏期间的改造^[22]。其中同生期的矿物转化作用是微孔形成的关键, 准同生期的白云石化则形成白云石晶间微孔, 重结晶作用和溶蚀作用等对微孔起积极改造作用。

图4 不同沉积环境中基质泥晶直径箱状图^[17]Fig. 4 Box plot showing the diameters of aphanocrystallines in matrix in different sedimentary settings^[17]图5 不同沉积岩相微孔率与总孔隙度交会图^[17]Fig. 5 Cross plot of microporosity and total porosity for different sedimentary lithofacies^[17]

2.2.1 矿物转化在微孔形成中起关键作用

矿物的转化作用是指同沉积期高镁方解石/文石质泥向低镁方解石泥晶转化,沉积期的暴露环境以及大气淡水形成的潜水透镜体促进矿物的转化,形成微孔^[22, 27, 29]。形状较好、物性较好且发育微孔的泥晶一般都是靠近层序界面之下^[17, 21, 88]。

从文石、高镁方解石到低镁方解石的转化是一个事实^[26, 64, 71-72, 78, 89-91],且不论是文石还是高镁方解石,转化形成的低镁方解石泥晶与原生低镁方解石泥晶无明显差别^[24, 29, 69, 92-93]。矿物成分的转化作用有利于晶间微孔的排列,是微孔形成的一个重要成岩作用^[2, 16, 20, 73-75]。前人对转化方式进行研究,提出了多种转化方式,包括晶粒生长作用^[89, 94]、新生变形作用^[95]、Ostwald熟化作用^[96]以及溶解-胶结作用^[2, 4, 16, 18-20, 71, 73, 97-98]。这些转化方式均指示:矿物成分的转化是在热力学的驱动下,文

石和高镁方解石组成的原生灰泥溶解,等量低镁方解石在空隙之间有序地形成泥晶。这个过程使得原本存在的不规则孔隙转化为均一排列的微孔。

2.2.2 早期白云石化作用收缩形成白云石晶间微孔

白云石化作用不仅有利于形成常规的孔隙,也有利于形成晶间微孔。白云石化作用包括交代白云石化作用和微生物诱导白云石化作用。

交代白云石化作用过程中,由于收缩作用会形成晶间孔和收缩缝^[99],特别是形成广泛的晶间微孔^[100],如在柴达木盆地英西地区湖相致密碳酸盐岩中,发现大量晶间微孔^[100-103]。在渤海湾盆地东营凹陷的白云岩中也观察到在颗粒内部形成晶间微孔^[76]。

在多个盆地、多个层位的微孔间都发现藻丝体等微生物的存在。例如在南海西沙地区西科1井中,可在扫描电镜下见到白云石晶间微孔中存在藻丝体^[104](图6a),在四川盆地高石梯地区龙王庙组的白云石晶间微孔也发现藻丝体(图6b),推断微生物可能与诱导白云石化而形成晶间微孔有关。另外,白云石化作用增强微孔的抗压实能力,使得孔隙在埋藏过程中的保存能力更强^[76]。

2.2.3 重结晶作用促进微孔形成及保存

大量的研究表明,经过泥晶化的颗粒更容易形成微孔^[12, 16, 20, 73, 105]。Bathurst提出的泥晶化是指原始结构颗粒转变为泥晶的过程,泥晶从较小的晶粒生长成较大晶粒,包括钻孔内沉淀无机泥晶碳酸盐岩,岩石内部文石的生物诱导沉淀以及骨骼重结晶。虽然这个过程与矿物转化作用相互混杂,但是该过程主要是微观结构上的变化,并主要由Ostwald熟化引导^[106-111]。另外一种重结晶是浅埋藏期,低镁方解石在Ostwald熟化影响下发生重结晶作用,使得低镁方解石泥晶排列更加均一,有利于晶间微孔的形成^[18, 69]。

2.2.4 溶蚀作用积极改造扩大微孔

溶蚀作用主要是扩大孔隙空间或者增大孔隙之间的连通性^[87],主要包括了同生期选择性溶蚀、埋藏期微生物热降解生酸溶蚀、油气生成和转化过程中的有机酸溶蚀、TSR(硫酸盐热化学还原反应)相关富含CO₂和H₂S流体溶蚀以及深层热液溶蚀^[112-114]。

同生期选择性溶蚀作用是在沉积物未固结成岩时,受到大气淡水淋滤对粒径较小的生屑颗粒内部进行溶蚀作用,形成铸模微孔^[23]。

在埋藏环境中,微生物热降解产生有机酸和CO₂,

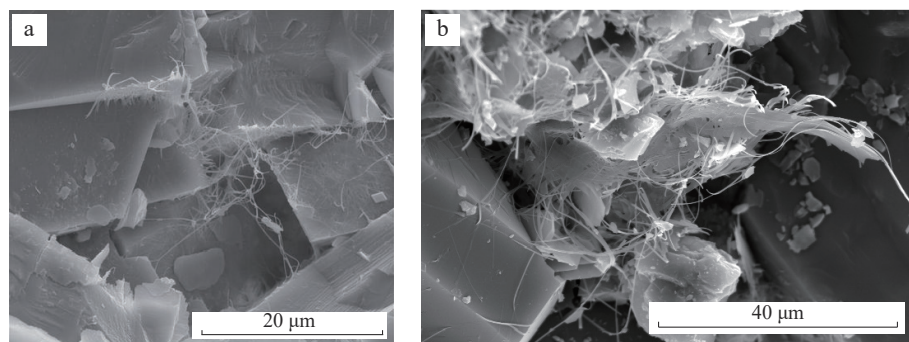


图6 白云岩晶间微孔中存在的藻丝体SEM照片

Fig. 6 Algal filaments in intercrystalline micropores of dolomites

a. 南海西沙地区, 三亚组二段, 白云岩中见微生物藻丝体, 埋深 1 255.92 m^[104]; b. 四川盆地, 龙王庙组, 白云石晶体表面附着丝状物, 埋深 4 486.49 m

产生的溶蚀作用不仅有利于微孔的形成, 而且可以抑制胶结物的形成, 有利于孔隙的保存^[31, 115-118]。真菌和细菌等组分的腐烂, 也会形成微孔^[119-120]。深埋藏期, 由于受到由烃类成熟、降解而产生的CO₂和有机酸^[121]等酸性埋藏流体的改造, 在深层条件下也可以有较高的孔隙度^[122], 在塔里木盆地顺南地区中-下奥陶统的实例中, 提出生烃规模也可能影响孔隙的发育方式, 生烃强度低可能导致孔隙仅发育于颗粒内部而非颗粒间^[31, 42]。深部溶蚀作用也受岩浆热液作用影响^[123], 同时热液溶蚀作用还受到热液上升速度、温度变化和构造断层裂缝位置等因素的影响^[124-125]。

2.2.5 胶结作用选择性充填保留微孔

受到表面化学相关的动力学影响, 碳酸盐岩表现为微孔的选择性保留, 在大孔与小孔共存时, 表现为大孔隙先充填, 微小的孔喉不易被方解石等胶结物充填, 可以成为主要和有效的孔隙空间^[20, 112, 126], 也使得微孔率得到有效提高。

2.3 碳酸盐岩微孔典型发育模式

微孔的发育与成岩作用密不可分, 其中矿物的转化作用以及埋藏期溶蚀作用的重要性得到广泛的认可。基于上述沉积条件以及成岩作用的调研, 认为文石海时期的低能沉积环境更有利于发育碳酸盐岩微孔, 而成岩作用则起关键作用, 并依据成岩序列建立理想条件下的微孔发育模式(图7)。

当沉积海洋环境为文石海环境时, 海水镁/钙比高, 碳酸盐岩的原生沉积物以高镁方解石和文石为主。由于海平面短暂下降, 低沉积地貌区域的沉积物暴露接受大气淡水淋滤, 在靠近海平面附近形成潜水透镜体, 高镁方解石和文石质的灰泥/颗粒在“Ostwald 熟化”作用的推动下, 先溶解再沉淀结晶, 形成排列整齐、

形状相似的低镁方解石泥晶(具有开放的骨骼格架或经具泥晶-粉晶结构的颗粒, 更容易实现矿物成分的转化作用)。在这个过程中, 早期存在的孔隙得到调整, 从不规则的孔隙转化为晶间微孔。进入浅埋藏期后, 受到 Ostwald 熟化作用的引导, 低镁方解石泥晶进一步重结晶形成更多晶间微孔; 基质中的一些微型藻类、真菌和细菌等微组分, 由于埋深增大、温度升高开始热降解, 形成溶蚀微孔; 烃类物质成熟产生有机酸和CO₂, 以及与 TSR 相关的富含CO₂和H₂S的流体开始对微孔进行溶蚀改造, 并使得微孔保存。

晚埋藏期, 构造裂缝为热液的流动提供通道, 一方面受到腐蚀性流体和岩浆热液的改造, 微孔溶蚀扩大, 改善了孔隙空间; 另一方面受到饱和流体的影响, 在孔隙之间形成胶结物, 特别是大孔的胶结充填, 导致微孔率上升。

3 微孔对岩石物性的影响

碳酸盐岩微孔的存在对岩石的物理性质, 包括岩石的孔渗特征、声波响应以及电阻率等有重要的影响^[17, 19, 127]。在以颗粒为主的碳酸盐岩中, 宏观孔隙控制岩石物理性质, 颗粒的大小对岩石的物理性质起主要的控制作用。而在微孔碳酸盐岩中, 微孔构成岩石中唯一的孔隙类型, 几乎所有的孔隙度都归因于微孔, 孔隙率在2%~35%, 渗透率在(0.01~100)×10⁻³ μm²^[1]。泥晶体形态和排列方式决定微孔的空间几何结构, 从而决定岩石的物理性质^[17, 19]。

3.1 孔隙度和孔喉半径

在石油开采中, 发现了部分灰岩具有高孔隙度、低渗透率的特征^[128-129], 称为微孔的“增孔”效应^[42]。微孔可以为碳酸盐岩储层提供大量孔体积^[130], 例如

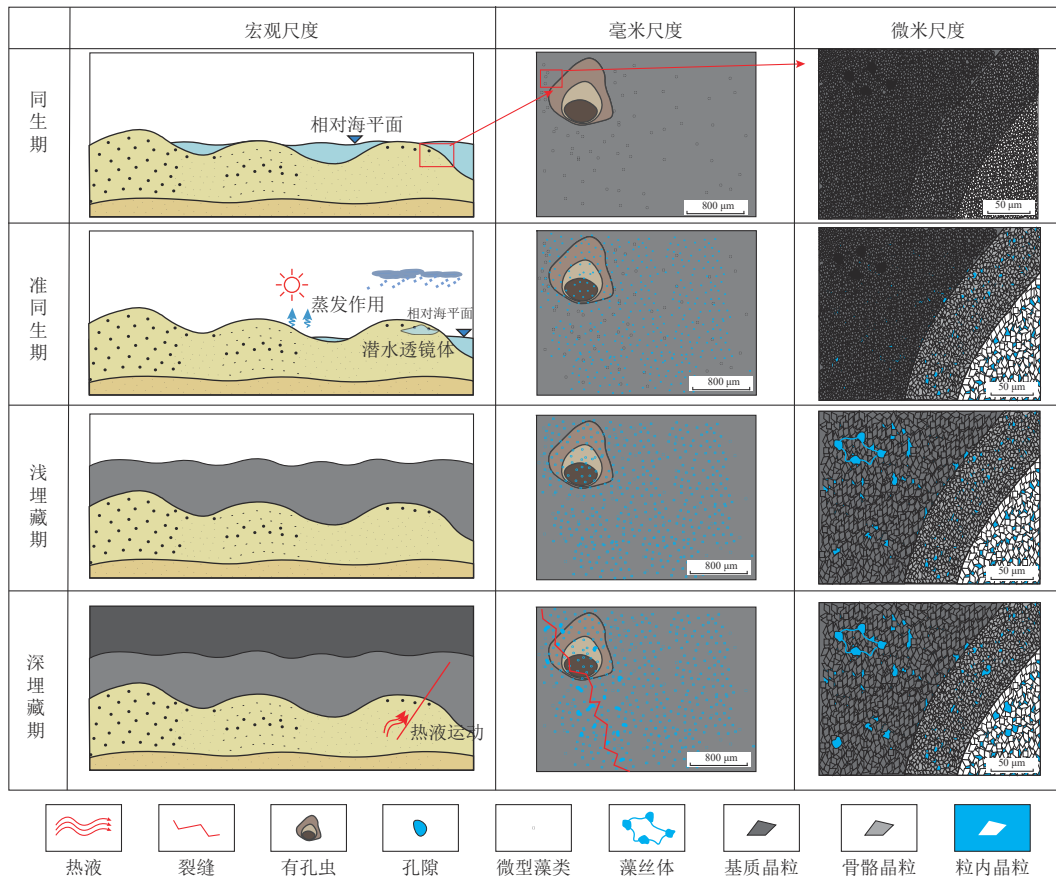


图7 典型碳酸盐岩微孔发育模式

Fig. 7 Typical development model of micropores in carbonate rocks

柴达木盆地英西地区微孔白云岩储集层最大孔隙度可达20%^[131]。通过对全世界多个微孔灰岩的调研发现微孔的孔隙度和孔喉半径与泥晶类型关系明显,其中半自形结构的微孔孔隙度和孔喉半径更大,自形结构的微孔次之,接合状结构的微孔最差^[17, 21, 88](图8)。

3.2 渗透率

微米-纳米级的微孔完全有能力赋存油气并允许其流动运移^[132-133]。一方面,微孔的发育在一定程度上可以改善储层的孔隙结构,起到连接孤立孔隙的作用,对储集性能的提高有很大的作用^[31]。通过实验证明,在总孔隙度一定的情况下,微孔率越高,渗透率越低^[13]。而另一方面,影响岩石渗透率的因素很多,包括泥晶的直径、孔隙类型以及泥晶的形态等等^[134]。例如,相同粒径的自形泥晶和接合状泥晶,渗透率范围变化明显;随着孔径分布范围减小或者是微孔之间的连通度增大,都会导致流体速度更加集中等^[17]。对于微米-纳米尺度的孔隙空间,渗透率不能完全表征

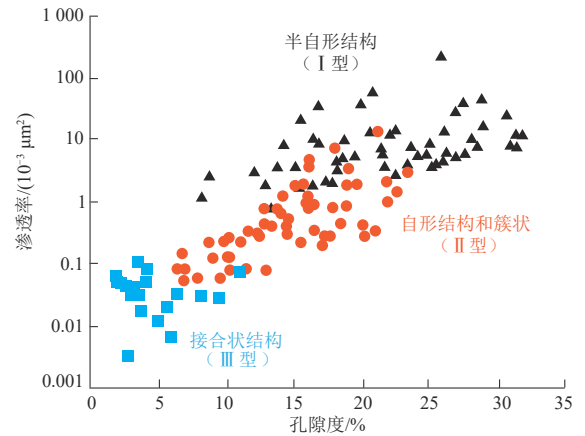


图8 以微孔为主的灰岩的孔隙度-渗透率对数关系曲线^[1]

Fig. 8 Relationship between porosity and permeability of limestones dominated by micropores^[1]

岩石的渗透性,邹才能提出孔隙连通率,同样能够对微米-纳米级孔隙空间的渗透性进行描述^[135]。实验表明影响渗透率的因素依次为裂缝、孔隙连通性、孔喉半径^[42]。

3.3 微孔对其他物性的影响

碳酸盐岩中的微孔还对岩石的电阻率、声学 and 力学等物理特性产生影响。微孔的存在使得电阻率明显下降:当储集空间主要为细微孔喉时,受导电迂回曲度的影响,在相同饱和度条件下,微孔型储层电阻率更低^[136]。另外,在研究中发现地层因素(饱和盐水电阻率与纯地层的电阻率之比)随孔隙度增加而减小,且泥晶结构类型与地层因子的分布关系密切^[88]。

微孔的存在使得声波速度减小,在孔隙度相同的情况下,微孔率越高,声波速度越小^[137](图9)。而泥晶直径和结构同样对声波速度产生影响,通过建立声波速度与孔隙度交会图,表明泥晶的直径减小以及半自形结构将会导致波速减小^[88](图10)。

另外,微孔有效地增加了岩石体积压缩系数,使得岩石受到应力后主要发生弹性形变,骨架刚度变

大^[138-139]。在实例的岩石压缩实验中,可观察到微孔碳酸盐岩的骨架体积弹性模量几乎不变^[140]。

4 研究趋势

一方面,原位微区的沉积、成岩环境定性-定量分析技术不断更新,原位微区元素、团簇同位素、U-Pb同位素测年等技术的发展,结合激光剥蚀和高精度质谱仪应用^[141],使得恢复沉积成岩的环境和过程更加准确,对明确微孔成因有重要意义。另一方面,微孔表征技术在进一步发展。新技术相较于传统的扫描电镜和显微CT扫描,能够更精确、更全面地表征微孔,如太赫兹成像技术(terahertz imaging)^[142]、激光扫描共聚焦显微镜(CSLM)^[143]以及扫描电镜大视域拼接技术^[144]等。其中激光扫描共聚焦显微镜不仅精度更大(2 μm /像素),而且能够识别微孔与宏孔的连接,评估孔隙的连通性,更有效对碳酸盐岩微孔进行定性和定量分析。

在不同研究目的和需求下,微孔表征技术方法在不断优化。①讨论微孔在包含大孔或裂缝的复合型孔隙系统中的贡献,通过数字模型的建立去讨论微孔的动态和静态特性,包括多物理场仿真、数字岩心等技术,如何使得数字模型更加真实地反应出流体在复合型孔隙系统中的流动特性是值得讨论的问题。②微孔碳酸盐岩孔、渗明显低于常规碳酸盐岩储层,需要建立一个标准去确定微孔碳酸盐岩的物性下限。

目前,碎屑岩微孔的物性下限研究已较为丰富,如通过驱替实验中压差饱和度关系^[145]、驱替启动压力与物性关系拐点^[146]、两倍水膜厚度与油气分子集合体之和^[147]、扫描电镜实测^[148]、流体性质受力^[133, 149]等方法研究和确定孔喉的下限。但以上方法受到样品数量和区域差异等因素影响较大,研究致密碳酸盐岩储层物性下限的方法要考虑流体渗流过程、孔隙结构、界面性质、流体性质以及温压条件等^[56]。

5 结论

1) 碳酸盐岩微孔是孔径小于50 μm 的微米-纳米级的孔隙,根据不同的研究目的有多种分类方案,以Kaczmarek的划分方案应用较为广泛,将泥晶形态和物性相结合,微孔分为I型、II型和III型等3种,分别发育粒状半自形、粒状自形或疏松簇状或紧密簇状以及部分接合状或紧密接合状结构,物性依次变差。微孔的表征方法主要分为流体侵入法、图像分析法和数学统计法,各类表征方法通过定性-定量组合实现对微

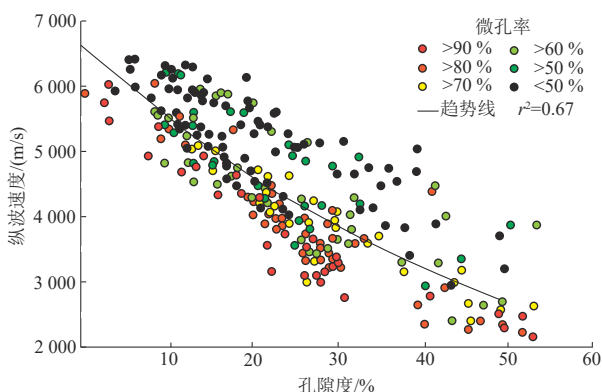


图9 微孔率和声速之间的反比关系^[137]

Fig. 9 Cross plot showing the inverse relationship between microporosity and P-wave velocity^[137]

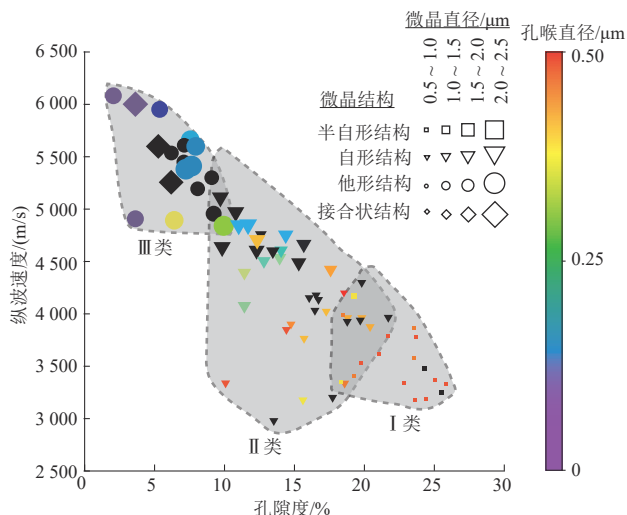


图10 总孔隙度与纵波速度交会图^[88]

Fig. 10 Cross plot of total porosity and P-wave velocity^[88]

孔的综合表征。

2) 微孔主要发育于碳酸盐岩的基质,其形成受到矿物、沉积条件以及成岩作用的控制。原生矿物成分是高镁方解石和文石以及颗粒具有开放原生微观结构的更容易发育微孔。海水环境为文石海、受大气的淡水影响且水体能量低的沉积相更容易发育微孔。同生期的矿物转化作用、白云石化作用是微孔形成的主要成岩作用,重结晶作用、溶蚀作用以及微组分热解生烃等对微孔起建设作用。

3) 微孔类型、孔径、泥晶形态、微孔连通性和孔喉半径等特征对微孔的物性有不同程度的影响。泥晶结晶程度越高,泥晶直径越小,孔喉半径越大,导致孔隙度和渗透率越高,电阻率越低,声波波速越小。

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